# APPENDIX NO. 7

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Anima .

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## MICROPALEONTOLOGICAL REPORT

STRATIGRAPHY

of the

FORAMINIFERAL SEQUENCE

in

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## SELENE # 1,

#### GIPPSLAND BASIN.

## for: PHILLIPS AUSTRALIAN OIL COMPANY.

March 3rd, 1983.

David Taylor, 23 Ballast Point Road, Birchgrove, 2041. AUSTRALIA. (02) 82 5643. SELENE # 1

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# STRATIGRAPHIC SUMMARY.

	Sample				
	Depth			ISTRAT UNIT &	E-LOG
		ZONE*	AGE†	PALEOENVIRONMENT¶	· PICK
	1270	B-2	LATE to	CANYON FILL MEMBER OF GIPPSLAND	
	to	to	MID	LIMESTONE shelf edge at base ( $\simeq 200$ m	a)
	1715	c	MIOCENE	filling to shallower depth (<200m)	•
		C	HLUCL	at top	
			<u> </u>		1740 or 1810
	1815	poor		CANYON FILL MEMBER of GIPPSLAND	
	1815 to	assemblages	MID	LIMESTONE. Upper Slope situation	
1	2232	? D-1	MIOCENE	(400-200m). High energy with	
-	dia are un en	• -		slumping at base	
		- 1			2234
1	2234	D-1	MID	TASMAN SEA CARBONATES with deep	
-	to	to	MIOCENE	water oozes. Anoxic between	
	2400	D-2	· · · · · · · · · · · · · · · · · · ·	2275 & 2234. Oxic between 2800	
•	2500	F	EARLY	& 2300 where energy regime was	
	to 2800	to N-1	MIOCENE	higher with influxes of detrital	
	2800	H-1		within the set of the	2807
ww		m.y.)~~~~~	wwwwwwww		NVVV ZOUI
	2810	- •	EARLIEST	LAKE ENTRANCE MARL equivalent	
	to	J-2	OLIGOCENE	shelfal environment - high	
-	2820			energy with detrital quartz etc.	
1				Paleodepth increasing upsection	
-				to 100-200m.	2821
		<u></u>			2021
ı	2822	K	LATEST	COLQUHOUN FORMATION	
-			EOCENE	Intertidal (<10m)	
$\sim$		m.v.) vvvvv	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~		VVVV 2822.5
-	2826	··· 2 - ·	late	GURNARD FORMATION	
1	to	N	MID	"Greensands" Intertidal near	
-	2838	· · · · · · · ·	EOCENE	entrance of Estuarine system.	
$\sim$		m.v.) vvvvv			vvvv2839
	2840	··· · ·	early		
	to	ο	MID	FLOUNDER	,
	2842		EOCENE		
?	? ?	?	?	FORMATION	2842
	2843	low	EARLY	equivalents Barrier/Dune/	
	to	diversity		lagoonal system.	
-	2852	- · ·	dinoflagella		
	2848	no	EARLY		
-	to	planktonics			
-	2852	•	dinoflagella	ates	
ww	····· (≃17	7 m.y.) vvvv		~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	vvv?2870
	2879.5	?	Latest	Swamp/Marsh ponds slightly	
	to	no	CRETACEOUS	saline at 3020 with arenaceous	
-	3151	planktonics		foraminiferal assemblage.	
			pollen (no		
			dinoflagella	ates)	
-	+ Cumma	of results	of examinat	ion of fifty sidewall cores and one c	conventional
-		sample as lis			JOINT CITELE
-		-			• .
-				aminifera with available palynologica	
	See pa	alynology rep	ort by Helene	e Martin & David Taylor for final syr	nthesis.
r	Detai?	led planktoni	c foraminifer	ral distribution on Table 2 for Eocer	ne to base Mioc
·				e l gives reliability for planktonic f	foraminiferal
/ <b></b>		boundary pic			
- <b></b>	¶ Inter	prototions ba	and on distr	ibuiton of planktonic, benthonic and	nlant microfos
	as we	11 as other s	ediment grain	ns (>.075mm). On Table 3 for Eocene	to base Miocen
٦		5 for Miocen			<b>-</b>
1		• • • • • • • • •	•••		

#### **BIOSTRATIGRAPHY.**

This discussion is based on the distribution of Tertiary planktonic foraminifera in the Gippsland sector of the Tasman Sea Provence. At the time of compilation of this report, palynological results were not complete, but note is made of preliminary results where applicable. The palynological compilation report by Helene Martin and David Taylor will contain synthesis of the planktonic foraminiferal and palynostratigraphy, with any necessary reconciliation.

### LATE CRETACEOUS-? to 3181 to 2879.5m.

Age is based entirely on identification of the spore/pollen Zone of *Tricolpites longus* by Helene Martin. No planktonic foraminifera were found in four samples examined. However, the sidewall core at 3020m contained a late Cretaceous/Paleocene morphologically primitive arenaceous foraminiferal assemblage similar to those described from the Otway Basin by Taylor (1965).

#### EARLY EOCENE - 2870 to 2843m.

Age determination based on dinoflagellates (see Palynology Report), as there were only sporadic occurrences of low diversity assemblages of *Globigerina* spp which were ubiquitous in both early and mid Eocene times.

#### MID EOCENE ZONE O - 2842 to 2840m.

The association of *Globorotalia centralis*, *G. collactea*, *G. turgida* and *Globigerina frontosa* places the assemblage at 2842 in a biostratigraphic position equivalent to Zone Pll of Blow (1979); thus deposition took place in the early part of the Mid Eocene.

#### MID EOCENE ZONE N - 2838 to 2826.

At or near the top of the mid Eocene on range overlap of elements in these assemblages. In the Tasman region (refer Jenkins, 1974) *Globorotalia collactea*, *Globigerina primitiva* and *G. angiporoides* all became extinct at or near the top of the Mid Eocene; whilst *Globigerinatheka index* and *Globorotalia aculeata* first appear in mid Eocene and continued into the late Eocene. Middle Eocene dinoflagellates are recorded in this interval in Selene # 1. A hiatus with an estimated time span of 7 million years was evident at 2839m (E-log) between Zone 0 and the overlying Zone N assemblages. The apparent abbreviation of the Zone 0 interval may have been due to erosion during the onslaught of the Zone N marine ingression.

#### LATE EOCENE ZONE K - 2822.

After a hiatus of some 2 million years, sediment was deposited at the top of the Eocene. This very high Eocene placement is evident by the presence of the keeled morphotype of *Globorotalia cerroazulensis cocoaensis* (refer Stainforth et al, 1975, p.258). The very thin development of Zone K (1.5m on E-log) was followed without a discernible break, by earliest Oligocene, Zone J-2 assemblage at 2820m.

The planktonic assemblage at 2822 was surprisingly diverse with warm water elements, such as *G. cerroazulensis cocoaensis*, whilst the early Oligocene faunas were composed entirely of species endemic to the Southern Ocean (Jenkins, 1974), with a total absence of tropical elements. This faunal change from cosmopolitan to parochial reflected the rapid paleotemperature deterioration in the Southern Ocean on the Eo/Oligocene boundary (Shackleton & Kennett, 1975 and Loutit & Kennett, 1981, p.60). All evidence presented indicates strongly that the sample at 2822 was deposited 37 million years ago, on the Eo/Oligocene boundary.

#### EARLY OLIGOCENE ZONE J-2 - 2820 to 2810m.

Assemblages contain elements typical of the Tasman Early Oligocene Zone of *Globigerina brevis* (Jenkins, 1974). The widespread Oligocene Hiatus of Gippsland (the *COBIA EVENT* of Taylor, 1983) and in the Tasman-Coral Sea region (Kennett et al, 1975 and Loutit & Kennett, 1981, p.57), was apparent at 2807 (E-log pick) with a time gap of some 12 million years.

#### MIOCENE - 2800 to 1270.

Sedimentation resumed in earliest Miocene times, as is evident by the presence of Zone H-1 assemblages between 2800 and 2765m. Above 2800, a continuous sequence of Miocene deep water and canyon carbonate sediments were present with the highest sample examined (at 1270m) representing the Late Miocene Zone B-2. A sampling gap of 100m between 2500m and 2400m occurred over the levels of occurrence of Zones E-2 & E-1. Usually those E Zone faunas occupy a very thin sediment interval in Gippsland deep water sequences. Moreover, the Zones E-2

& E-l time span was very brief as species evolution was accelerated by oceanic warming.

Biostratigraphic control was poor to non-existant in the canyon carbonate fill sequence at and above 2232m. This is a common phenomenon in Gippsland wells, due to the high energy regime which sorts and concentrates mainly very small sized specimens, which, being juvenile are specifically indeterminate. Another factor, regarding species recognition, is the poor preservation, cuased by carbonate diagenesis of the fill. It is assumed that Zone D-1 occupied the thick interval from 2232m to at least 1815m in Selene # 1.

#### PALEOENVIRONMENT and ROCK STRATIGRAPHY.

Some of the remarks below, may be expanded or slightly amended in the final paleoenvironmental and rock stratigraphic conclusions are drawn in the paleontological synthesis section of the Selene # 1 palynology report.

#### ? to 3181 to 2879m - Latest Cretaceous - Latrobe Delta Complex.

Predominantly non-marine, but with at least one marginal marine episode at 3020m; evidenced by an association of the arenaceous foraminifera *Haplophragmoides* spp. with pellet glauconite and biogenic pyrite. Such a benthonic fauna was euryhaline, withstanding fluctuation in salinity to as low as  $4^{\circ/}$ oo as well as anaerobic conditions (Taylor, 1965). Dinoflagellates were not recorded in this sample which was usually the case in sediment containing *Haplophragmoides* spp assemblages from the Late Cretaceous to early Tertiary of the Otway Basin.

2870 to 2840m - EARLY ECCENE to MID ECCENE - FLOUNDER FORMATION EQUIVALENTS. In this interval, the dominant sediment grains (>.075mm) are frosted, pitted and/or impact fractured quartz grains, features probably cuased by eolian processes (Margolis & Krinsley, 1974). Limonitic clays are common and could have been from paleo-soil horizons. Within this interval, were indications of marine ingressions into marginal marine environments. Dinoflagellates occurred throughout, with sporadic planktonic and arenaceous benthonic foraminiferal associations in five out of the nine sidewall cores. Foraminiferal frequency increased upsection, reaching a peak at 2842m (refer Table 5). Distribution of glauconite and biogenic pyrite also demonstrates an up-sequence frequency increase. One sample, at 2845m, contained an

appreciable amount of crystalline carbonate, which was probably dolomite, though possibly siderite. The total of these observations is that sedimentation took place within a barrier/dune/estuarine regime analagous to that of the present day Gippsland Lakes - Ninety Mile Beach system (Taylor, 1983).

This interval is equated with the Flounder Formation having been deposited over the same time-span and within a marginal-marine regime. However, the sedimentary facies in Selene differs from that defined for the Flounder Formation. But the rock type is dependent on the exact position at any one time that sedimentation took place within such a barrier/dune/estuarine system. The typical siltstones and mudstones of the Flounder Formation were deposited in deep protected estuaries behind and to the lee of the dunes, whilst the Selene sequence was deposited on or just windward of the dunes. It should be noted that the lower part of the Formation in the type section contains an abundance of quartz with at least one coarse sandy lens (refer Flounder # 1 tabulation in Taylor, 1983, p.13).

#### 2838 to 2826m - LATEST MID EOCENE - GURNARD FORMATION.

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Sedimentation took place in a lagoonal situation in proximity to the entrance to the system from a shallow platform continental shelf. The basal sample at 2838m was the only one to contain quartz grains, sculptured by eolian processes. No doubt this was a surface of the underlying unit, reworked during the onslaught of the late Mid Eocene ingression. Apparent abbreviation has already been noted at the top of the Flounder Formation in Selene (refer p.2 this report).

Age, faunal characteristics and sediment types are all consistant with a designation of Gurnard Formation for this unit.

#### 2822 to 2810m - LATE EOCENE to EARLY OLIGOCENE - COLQUHOUN FORMATION and LAKES ENTRANCE MARL EQUIVALENTS.

A sequence of fine quartz sandy marls were deposited at the very top of the Eocene and continue into early Oligocene times, before the effects of the COBIA EVENT resulted in all the mid to late Oligocene sediment being absent in this sequence. The encroachment of the late Eocene-early Oligocene transgression is demonstrated by changes in the benthonic foraminiferal assemblages and the high percentages of planktonic specimens in the total

foraminiferal faunas. Rapid increases in paleodepth are evident with a litteral situation in the latest Eocene progressively becoming an outer shelf one in the early Oligocene (refer Table 3).

On lithology alone it is not possible to differentiate between the Colquhoun Formation and the Lakes Entrance Marl equivalents. However there are marked E-log character changes at 2821m which correspond to the paleoenvironmental changes from the Late Eocene litteral fauna at 2822 to the Early Oligocene shelfal faunas from 2820 to 2810m. Therefore:-The E-log interval 2822.5 to 2821m is considered to be the Colquhoun Formation:

whilst the E-log interval 2821m to 2807m is Lakes Entrance Marl equivalent.

#### 2800 to 2234 - MIOCENE - TASMAN SEA CARBONATES and OOZES.

A thick sequence of biogenic carbonates with fluctuating proportions of quartz grains and non carbonate silts and clays. There is no meaningful pattern to variations in the degree of carbonate diagenesis (refer Table 5).

The entire unit was deposited on the upper part of the continental shelf in estimated water depth between 200 and 400m. So there was a marked environmental disruption associated with the *Cobia Event*, as the early Oligocene deposition was on the continental shelf.

A change in available oxygen to the depositional surface is noted between 2275 and 2245m; conditions at and below 2275 were oxic, whilst those above 2275 were distinctly anoxic. The anoxic conditions are evident from the sudden appearance of 30% biogenic pyrite in the sediment grain spread. Bioturbation, with faecal pellets and worm tubes accompanied this incoming of pyrite. Benthonic foraminiferal assemblages also reflect a physico-chemical change; especially with the presence of a calcareous spiculitic shelled form referable to the unusual genus *Carterina*. A similar change from aerobic to anaerobic deposition was documented in Helios # 1, where it occurred in Zone F, earlier than the Zone D-1 occurrence in Selene # 1.

## 2232 to 1270m - MIOCENE CANYON FILL MEMBER of the GIPPSLAND LIMESTONE.

The base of the canyon fill sequence is recognised at 2232 by the deterioration in frequency and diversity of foraminifera. The very high proportion of small sized specimens of planktonics resulted in the loss of biostratigraphic control from 2232 and 1815m (refer p.4). Progressive progradation and infilling of the canyon is interpreted from microfossil distribution. The change from an upper slope to shelf edge situation was marked by the influx of siliceous sponge spicules at 1715m and a general improvement in the variety and preservation of the foraminifera; biostratigraphic control was re-assumed at this level.

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# TABLE 1

# MICROPALEONTOLOGICAL DATA SHEET

ВА	s I	N: GIPP	SLAND				ELEVA	TION: KB	: _2	3m GL:	-268	m
WELL	NAI	AE: SELE	NE # 1				TOTAL	DEPTH:	<u> </u>	· · · · · · · · · · · · · · · · · · ·		<u> </u>
			HIG	ΗE	ST D	АТ	A	LO	WE	ST D	AT	
AG	; E	FORAM. ZONULES	Preferred Depth	Rtg	Alternate Depth	Rtg	Two Way Time	Preferred Depth	Rtg	Alternate Depth	Rtg	Two Way Time
PLEIS-		<sup>A</sup> 1				ļ						
ЦĞ		A <sub>2</sub>		<b></b>								
ι υ		A <sub>3</sub>				<u> </u>						
PLIO- CENE		A <sub>4</sub>									+	
	<u>ല</u>	B B									+	
	LATE	B <sub>2</sub> C	1270	1				1415	1			
	ш		1515	1	<u></u>			1715	1			
យ	- Ц	D D	2215	2	2234†	0		2240†	1		+	
и ы	D D	D <sub>2</sub> E <sub>1</sub>	2245	0				2400	0		+	
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о н	W		* *	-					0		+	
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EOC-	ana	Pre-K	2822	1		+		2842	1			
			2020	1	<u> </u>		1	<u>u</u>		·L		4
co	MMEI	NTS The ve	ery thin r	epre	sentation	of I	D-1, due	to advers	se er	nvironment	al	
		factor	<u>cs operati</u>	ng i	n Miocene	can	yon at &	above 223	32.	Actual t	:op	
		ma	ay be as h	igh	<u>as 1740m E</u>	<u>E-lo</u>	g	<u></u>				
			ing gap of		m, probabl	ly r	esponsib	le for nor	<u>1-rec</u>	cognition	of	
		·	E-1 & E-2									
		<u>I</u> Zone I	N (late Mi	d Eo								
		-	<u>e) at 2842</u>				ity Earl	y to Mid I	Eocei	<u>ne plankto</u>	<u>mic</u>	
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	ATI)							ge (high confident		).		
		2 : 3 :			<ul><li>Close to zo</li><li>Complete</li></ul>			able to interp	ret (la	ow confidence	:).	
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		<pre>indet - poor p frontosa senni linaperta centralis collactea turgida increbescens angiporoides m primitiva aculeata cerroazulensi angiporoides a gemma brevis tripartita munda nana continuosa tapurensis venezuelana woodi woodi bulloides lissimilis lissimili lis</pre>			
	SIDEWALL CORES CONVENTIONAL C Depth in metre	<pre>indet - poo frontosa senni senni linaperta a centralis a collactea a turgida a culactea angiporoide primitiva ka index a aculeata a cerroazule a gemma brevis tripartita a munda a continuosa tripartita a nana brevis tripartita a nana a continuosa tapurensis venezuelana praebulloides dissimilis dissimilis a siakensis</pre>		Donth	
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• ]	2843.0-		?		to EARLY
	2845.0→ 2848.0→		?		EOCENE
•	2852.0. 2855.0.	0 0	Ļ	-2852	
	2860.5		?		?
<b>I</b> -	2870.0→ 2875.0→		ļ		
	2933.0+	<b>15</b>	2		2
	3020.0→ 3121.0→		İ		
	3151.5. CC # 1				
					1
]	KEY:	<pre>° = &lt;20 specimens N.F.F. = no foraminifera foraminif</pre>	ounđ		
		D = Dominant >60% specimens? = hiatus - uncertai	n		
		? = determination queried			
	T	ABLE 2: EOCENE to EARLY MIOCENE PLANKTONIC FORM	MINIFE	ERAL	
		DISTRIBUTION - SELENE # 1.			
	refe	r Table 4 for Miocene Distribution above 2800.0m.			
	Davi	d Taylor, 18/2/83.			

BENJ	HONIC FORAMINIFERA	RESIDUE	LITHOLOGY		1PALEO- ENVIRONMEN	TAL			
		MAJOR COMPONENTS	MINOR COMPONEN	NTS	ASSESSMEN	T			
Taylor sis	a ormis mois eous) eous)	<pre>Y: recrystallised biomicrite o: m-c ang. qtz. t.t.f. qtz. sandy marl</pre>	frosted peilets		tem Om) Om) (<40m)	CHANGE	FOR	LANKTO AMINIF TRATIC	ERAL
SIDEWALL CORES & CONVENTIONAL CORES Depth in metres Haplophragmoides sp.B T H. cotundata H. cf. incisa Ammodiscus parri Bathysiphon angleseaens Eathysiphon angleseaens Cothoides thiara C. perforatus	Nuttalides cf. trumpy Nuttalides cf. trumpy Bulimina bortonica Gyroidina subrealandica Gudyrina convera Saibboina australis Gadyrina convera Textularia subgriobsa Taeeisierella tartilarformis Sipboutgerina canarlensis Oridoralis umbonatus Sipaeroidina bulloides Bathysiphon (porcelaineous) Bathysiphon (porcelaineous) Discammina sp. "Cyclammina globiginiformis	<ul> <li>o: polymodal quartz silty sandstone</li> <li>G: pallet glauc</li> <li>AV: frosted £ fractured quartz</li> <li>P: pyrite</li> <li>: qtz. sandy</li> <li>siltstone</li> <li>\$</li></ul>	r tlet	foram count	ne/marsh sys ne/marsh sys antrance (<1 inner shelf (<100m) f (<200m)	E-LOG CHARACTER	ZONE	Depth at Base	AGE
2800.0.		*****	YAA A	200 9		1	H-1.	2800	EARLY
2810.0, 2815.0, 2818.0, 2818.0, 2820.0, *	· · · · · · · · · · · · · · · · · · ·	+,+,+,+,+,+,+,+,+,+,+,+,+,+,+,+,+,+,+,	A TTA	250 250 7 250 8	5	2821	J-2	2820	EARLY OLIGO- CENE
2822.0, 2826.0, 2825.0, 2823.0, * x * x * *	••••	+.+.+.+.+.+.+.+.+.+. 		-100 9 -100 6 100 6	nfmmptim	.5	N	-2822	LATE EOCENE 1atest MID
2838.0, x • 2840.0, 2842.0, •		VAVAVAVAVAVAVA PGGGG		7 7 7 7 25 8	nfinipanni	+2835 2839 2843		-2838	EOCENE MID EOCENE
2843.0, 7 7 7 2845.0, N.F.F. 2845.0, N.F.F. 2852.0, 7 7 7 *	· · · · · · · · · · · · · · · · · · ·	7 7 7 7 7 7 7 7 7 7 7 7 7 7	??? λλ λ	7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7		2847 2847 to 2850	-? ? ?	-2842 -2852	to EARLY EOCENE
2855.0. 2660.5. 2870.0. 2875.0. 2933.0.		ΔΔΥΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑΔΑ	A A A A A A A A A A A A A A A A A A A	20 7			?		7
3020.0+ * * * 3121.0+ } 3151.5+ }N.F.F		0+.000000- 		10 -			?		?
<pre>KEY: * = &lt;20 specimens x = &gt;20 specimens D = Dominant &gt;60% specimens</pre>	N.F.F. = no foraminife		A = abundant 1-5% gra C = common >20 grains r = rare <20 grains		TPaleowater de in parenthese		4	L	<u>L</u>

D = Dominant >60% specimens

? = determination queried

TABLE 3: PALEOENVIRONMENTS - EOCENE to EARLY MIOCENE - SELENE # 1

Refer Table 5 for Miocene Paleoenvironments (2800-1270m)

David Taylor, 22/2/1983.

	DE NUZONIC BODININIEEDI		
	PLANKTONIC FORAMINIFERA	BIOSTRAT	TGRAPHY
	L.) (<.2mm) S.) S.) S.) S.)	, ,	
SIDEWALL CORES Depth in metres	<pre>G'ina woodi connecta G'ina woodi woodi G'ina bulloides Cat. dissimilis Cat. dissimilis G'quad dehiscens (S. G'quad dehiscens (S. G'quad altispira G'quad altispira G'quad altispira G'quad altispira G'quad altispira G'quad altispira G'quad altispira G'alia siakensis/maye G'alia siakensis/maye G'quad dehiscens (S. G'alia siakensis/maye G'alia praendica (S. G'alia praenenardii G'alia praenenardii G'alia praenenardii G'alia praenenardii G'alia praenenardii G'alia praenenardii G'alia praenenardii G'alia praenenardii G'alia praenenardii G'alia anda G'alia anda G'alia scitula G'alia scitula G'ina nepenthes G'ina decoraperta G'alia acostaensis</pre>	zonE Depth at Base	AGE
1270.0 <sub>→</sub> 1315.0 <sub>→</sub>	x x D x x x x x x x x x x x x x x x x x	в-2	LATE MIOCENE
1415.0 <sub>→</sub> 1515.0 <sub>→</sub> 1615.0 <sub>→</sub> 1715.0 <sub>→</sub>	x x x x x x x x x x x x x x x x x x x	C 1715	
$1815.0_{\rightarrow}$ $1915.0_{\rightarrow}$ $2015.0_{\rightarrow}$ $2115.0_{\rightarrow}$		? 2115	MID
2215.0 <sub>→</sub> 2225.0 <sub>→</sub> 2229.0 <sub>→</sub>			MIOCENE
$2232.0_{\rightarrow}$ $2234.0_{\rightarrow}$ $2237.0_{\rightarrow}$ $2240.0_{\rightarrow}$	xx x ° ° x°°x xx ° ° x°°° xx ° ° x°°	D-1 2240	
2245.0 <sub>→</sub> 2275.0 <sub>→</sub> 2300.0 <sub>→</sub>	x x     x x x     x x x x x     x x x x x       x x     x x     x x x x x x     x x x x x       x x     x x     x x x x x x     x x x x       x x     x x     x x x x x x x x x x     x x	D-2	
2400.0, 2500.0, 2635.0, 2705.0,	x * * * * * * * * * * * * * * * * * * *	F 2500 G	EARLY
$2725.0_{\rightarrow}$ $2745.0_{\rightarrow}$ $2765.0_{\rightarrow}$ $2785.0_{\rightarrow}$	xxxx °°D x xx °°D x xx ° D°°°°° D		MIOCENE
2800.0	x x x ° ° ° ° x ° Mathematical Structure Stru	J-2	EARLY OLIGO- CENE
;	<pre>= &lt;20 specimens</pre>	- <b></b>	

TABLE 4: MIOCENE PLANKTONIC FORAMINIFERAL DISTRIBUTION - SELENE # 1.

David Taylor, 25/2/1983.

	BENTHONIC FORAMINIFERA	RESI	DUE LITHOLOGY			1PALEO- ENVIRONME	NTAT.						
		MAJOR COMPONENTS	MINOR COMPON	VENTS		ASSESSME					i.		
RES	riniformis aineous) cosa is cosa is corris c	Sp=siliceous sponge spicules A=recrystallised calcarenite -	aecal pellets			(<100m) on (<200m) on (≈200m) (00m) )-400m)	TER CHANGE	BIO	STRATI	GRAPHY		·	
CONVENTIONAL CORES Depth in metres	Discammina sp. "cyclammina "incisa Trochammina gibbigerif Frochammina gibbigerif Frochammina gibbigerif Bachysiphon (Porcelais Cibicides temperate & Nodosaria spp. Grodoralis umbonatus Cibicides temperate & Nodosaria spp. Bracinisiphon sp. (Jagena spp. Eracinisiphon sp. Nodosaria spp. Castelundina zealandica (Dibicides karteriform Marthoticila corniunti Sphaeroidina bulloide: Sphaeroidina bulloide: Cratcobulinta sp. (with 7 Carterina sp. Cibicides seudounger Cibicides seudounger Cibicides seudounger Cibicides subhaidinge Cibicides subhaidinge Cibicides subhaidinge Spirillina sp.	A=silty calcarenite m=biogenic micrite	glauconite pellets pyrite limonitic clay f-c ang-subrd qir fish fragments sponge spicules 7 worm tubes f 7 a bryozoal fragments echinoid spines	foram count	tonic s	Mid-Inner Shelf (<100m Outer Shelf Canyon (<2 Shelf Edge Canyon (*20 Slope Canyon (*200m) Upper Slope (200-400m)	MAJOR E-LOG CHARACTER	ZONE	Depth at Base	AGE			
70.0.+ 15.0.+ 15.0.+ 15.0.+	• x • x • x x x • • • • • • • • • • • •	ΛΠΩ ΠΟΠΟΛΟΛΟΛΟ Ο ΛΟΛΟΛΟΛΟΛΟ Ο ΠΟΛΟΛΟΛΟΛΟΛΟ Ο ΛΟΛΟΛΟΛΟΛΟ Σερία	А А А А С А	200 1000 500	75 90 95 95		21374	B-2	-1415	LATE MIOCENE			
$15.0_{+}$ $15.0_{+}$ $15.0_{+}$ $15.0_{+}$	x x x x x x x * * * * * * * * * * *	<u>,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,</u>	тгАА А	1000 1000 1000	90 98 90 98 98 98		1740 -or 1810	C	_1715	· · · · · · · · · · · · · · · · · · ·			
15.0. 15.0. 15.0. 25.0. 25.0.	x x x		λ	.? ? 100	?. ? 99 ?	SLUMP			_2115	MID MIOCENE			
32.0. 34.0. 37.0. +0.0.	• • • • • • • • • • • • • • • • • • •	AAAAAAAAAAAA Maanaaaaaaaaaa PPPPP Maanaaaaaaaaaaaaa PPPPP Maanaaaaaaaaaaaaaaaa PPPPP	<u>r</u> rλ λλ 	500 500	99 90 95 99	ANOXIC	2234	D-1	-2240				
45.0 <sub>+</sub> 75.0 <sub>+</sub> 00.0 <sub>+</sub> 00.0 <sub>+</sub> 00.0 <sub>+</sub>	x*************************************	nananananananan PPPP mananananananan PPPP manananananananananananan manananananan	AA CA	1000 1000 1000 1000	95 99 9 <u>9</u> 95	-2275-	~~~	D-2	_2400 _2500				
35.0 <sub>+</sub> 05.0 <sub>+</sub> 25.0 <sub>+</sub> 45.0 <sub>+</sub>	••x ••	илилияниянаянаянынынын ААААААААААААААА ААААААААААААААА	λ.	? 500 3000	99 7 99 98	OXIC		G	-2745	EARLY MIOCENE			
65.0+ 85.0+	• • x • • indet • • •		A r Ar A	.7 .200	80 7 98	h	2807	H-1	·····	EARLY OLIGO-			
x	<pre>refer Table 3 for Oligocene to Eocene palecenvironment data. = &lt;20 specimens = &gt;20 specimens = worn shallow water - displaced specimens.</pre>		A = 1-5% of grains C = >20 grains r = <20 grains	<b>i</b>		Paleowater in parenth		J-2		OLIGO- CENE			

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